

The First Find of the Early Jurassic Radiolarians in the Cherty Allochthons of the Samarka Accretionary Prism (Southern Sikhote Alin)

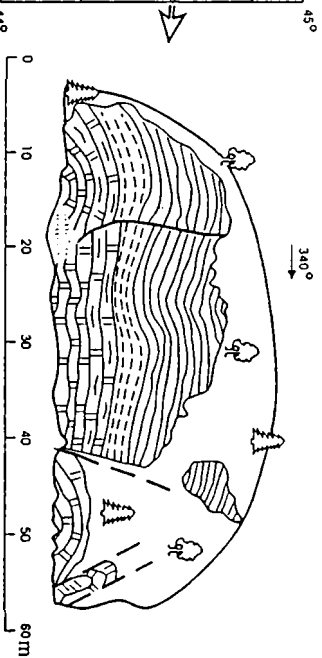
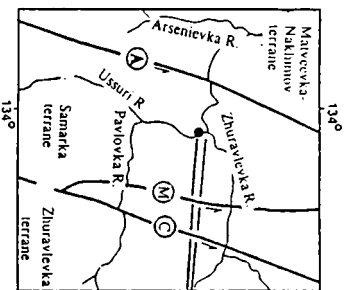
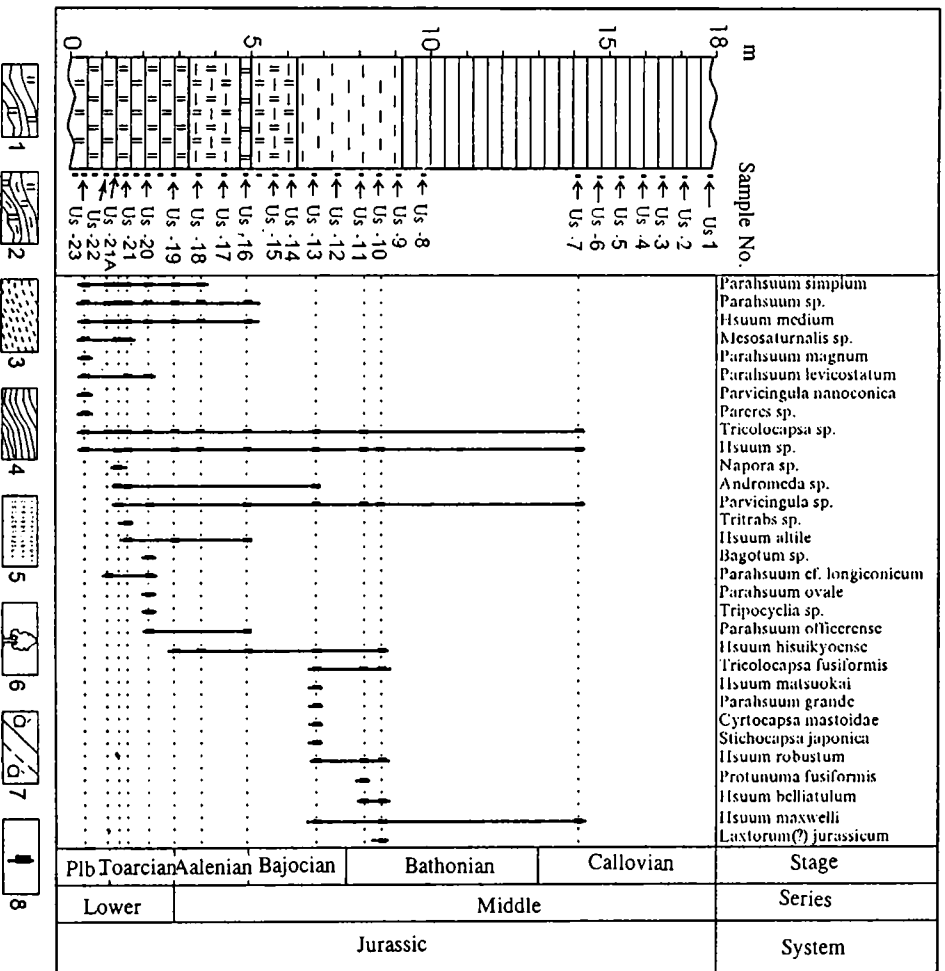
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New data on the finds of the Early Jurassic radiolarians in the cherty allochthons of the Samarka accretionary prism and the Middle Jurassic radiolarians (beginning with the Aalenian), in the turbidite matrix are given. The obtained data enable widening the chronostratigraphic range of the cherty allochthons (up to the Lower Jurassic inclusive) and to revise the date of inception of the Samarka accretionary prism. Photographs of the Early and Middle Jurassic radiolarians are supplemented.

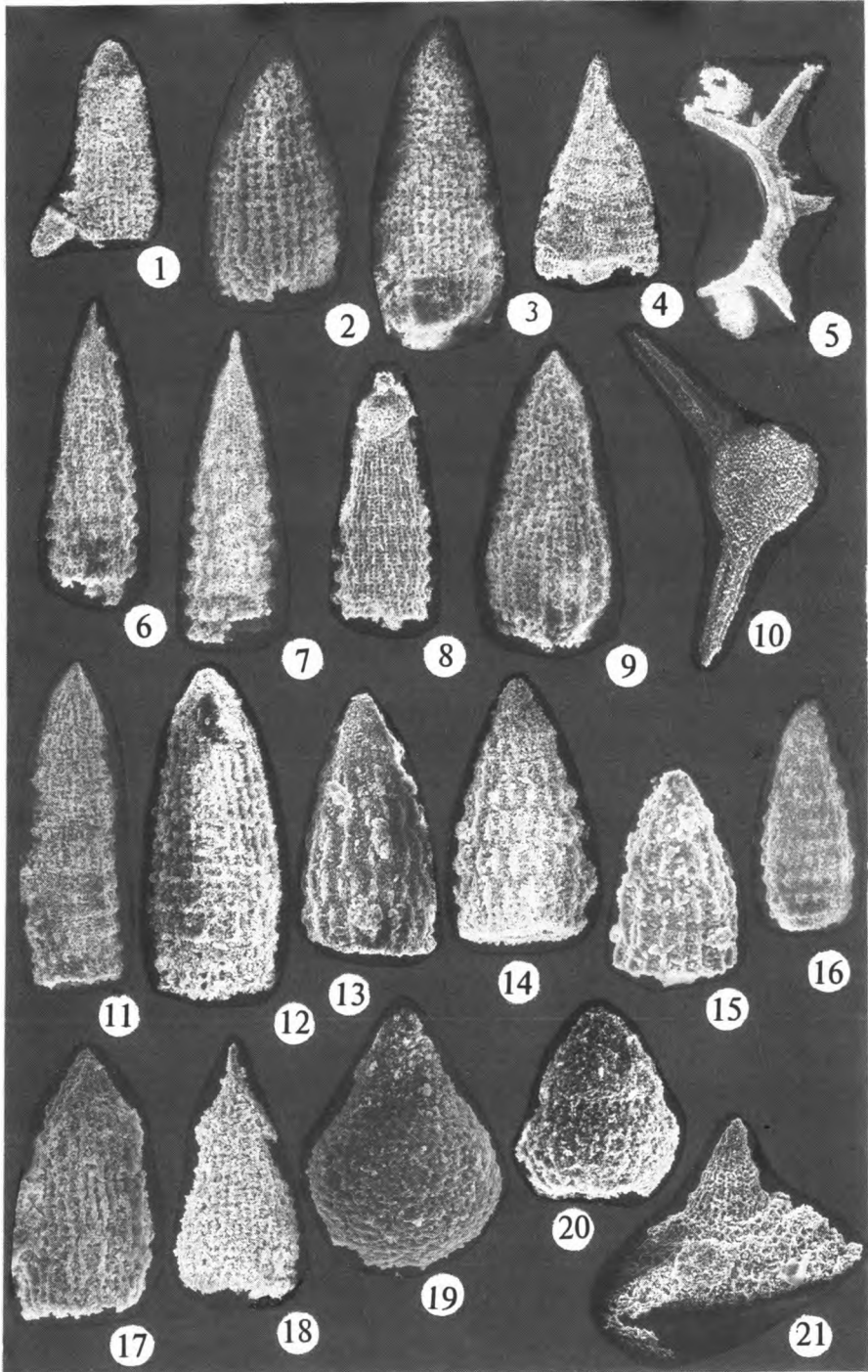
The Samarka terrane of the Sikhote Alin accretionary prism is a complex association of marginal continental and oceanic rocks [4], [6]. The marginal continental rocks are represented predominantly by distal arkosic turbidites and olistostromes; the oceanic, by reefal limestones and basalts that occur as blocks of various size in olistostrome (fragments of subaqueous highs and guyots) and ribbon cherts often associated with tholeiitic basalts that occur both as blocks in the olistostrome and as independent slabs with a length up to several tens of kilometers (fragments of abyssal plains). The previous isotopic datings of the cherts indicated their Late Permian and Triassic age [1], [4], [5]. The clastic rocks contain younger radiolarians of mostly Middle–Late Jurassic age [3], [5], [6]. And the relationships between the cherty and clastic sequences of the Samarka prism, differing in composition, age, and depositional environment, were insufficiently studied. At the same time, the solution to this problem is of great importance, firstly, for the structural interpretation of the accretionary prism and reconstruction of its evolution, and secondly, for the dating of subduction and accretion. In our opinion, the change from silicic to clastic sedimentation indicates that the fragment of an oceanic plate has approached the subduction zone. Therefore, the age of the last cherts and the first clastic rocks may indicate the date when this oceanic plate fragment started to subduct and, accordingly, the accretion process began.



The immediate contacts between the cherty and clastic strata are usually strongly deformed: quite common is the multiple repetition of these members resulting from both accretionary and the subsequent collisional tectonics. In some outcrops, however, we were lucky enough to observe gradual transition from the cherts to the overlying clastic rocks. Fig. 1 shows an example of well preserved transition discovered and studied in detail on the left bank of the Ussuri River 3.6 km upstream from the mouth of the Zhuravlevka River. This sequence is a fragment of the lowermost structural level of the Samarka prism formerly recognized as the Tudovak Formation [2]. At this level, stratified cherts are substantially predominant, less abundant are chert slabs, and the olistostrome horizons account for up to 5% of the total rock association volume.

The sequence consists of cherts (Fig. 1), the overlying cherty mudstones, and then silty mudstones and siltstones interbedded with sandstones. It should be noted that the cherts to clastics transition is gradual (by means of decrease in cherty components) within a 3 m-thick cherty mudstone sequence. We have sampled all lithologic rock varieties in search for microfossils. As a result, the following data were obtained: for the first time, the latest Early Jurassic radiolarians were identified in the cherts, and the Middle Jurassic radiolarians, in the clastic sequence (Fig. 1, Plate 1). A relatively rich radiolarian assemblage found in sample Us-23 from the lowermost part of the exposed cherty-clastic is characteristic of a wide chronostratigraphic range (Table 1). However, we determine its age as Pliensbachian, because the first appearance of *Hsuum medium*, *Parahsuum levicostatatum*, *Mesosaturnalis* sp., *Parvicingula nanoconica*, and *Parares* sp. is reported since the final Early Jurassic, most probably Pliensbachian–Toarcian [7], [11]. The upper age boundary was established on the basis that the radiolarians from the superjacent sample Us-22 are dated as latest Pliensbachian–Toarcian. Sample Us-22 contains a sparse radiolarian assemblage (Table 1) with most of forms characterized by a wide chronostratigraphic range. But the presence of *Parahsuum longiconicum* restricts the age to the interval between the latest Pliensbachian and earliest Toarcian [10]. Therefore, the Pliensbachian/Toarcian boundary lies in the immediate vicinity of sample Us-22. In samples Us-21A, Us-21, and Us-20, fairly rich radiolarian assemblages were identified with most of forms appearing late in the Early Jurassic and most abundant in the Middle Jurassic (Table 1). However, the presence of *Parahsuum longiconicum* restricts the upper age boundary to the final Toarcian. Considering that samples Us-21A, Us-21, and Us-20 were picked from higher stratigraphic levels than sample Us-22, their age can be defined

Fig. 1 Location map, outcrop sketch, and lithostratigraphic column of the transition (cherts to clastics) zone of the Samarka accretionary prism showing the radiolarian species distribution throughout the sequence. 1 – cherts; 2 – cherty mudstones; 3 – silty mudstones; 4 – turbidites; 5 – diluvial cone; 6 – soil and vegetation; 7 – faults: apparent (a) and inferred (b); 8 – localities of the radiolarian species in samples. Circled letters indicate faults: Arsenievka (A), Meridional (M), and Central Sikhote Alin (C).



as Toarcian. The radiolarian assemblage found in sample Us-19 shows a relatively narrow chronostratigraphic range – Toarcian–Bajocian, because it contains *Hsuum altile* [7]. At the same time, the presence of *Parahsuum simplum*, more characteristic of the Early Jurassic, restricts the chronostratigraphic range of this assemblage to the latest Toarcian–earliest Aalenian. Therefore, the Lower/Middle Jurassic boundary lies near sample Us-19 and coincides with the cherts to clastics transition.

Fairly abundant radiolarian assemblages characteristic mostly of the latest Early–latest Middle Jurassic were identified in samples Us-18 and Us-16 (Table 2) collected from the cherty mudstones of the transitional sequence between the cherts and clastics proper. The upper age boundary of these assemblages corresponds to the mid-Bajocian, because the superjacent sample Us-13 contains *Hsuum robustum* with a stratigraphic range between the middle and late Bajocian [9]. Samples Us-11 and Us-10 also carry rich Middle Jurassic radiolarian assemblages. Considering that the subjacent sample is dated middle–late Bajocian, the lower boundary of these assemblages is established as final Bajocian–earliest Bathonian. The upper limit corresponds to the final Bathonian, because the samples carry *Protunuma fusiformis* characteristic of the Bajocian–Bathonian [8]. The age of the last sample Us-7 is tentatively assumed as Callovian, because the subjacent samples are dated as Bathonian, and the sparse radiolarians identified in this sample are characteristic of the Middle to Late Jurassic (*Hsuum maxwelli*).

To summarize, 10 radiolarian assemblages have been identified within a 18 meter-thick continuous cherty-clastic sequence, characteristic of the following chronostratigraphic intervals, respectively: Pliensbachian, final Pliensbachian–earliest Toarcian, Toarcian, final Toarcian–earliest Aalenian, Aalenian, final Aalenian–earliest Bajocian, middle Bajocian–final Bajocian, final Bajocian–earliest Bathonian, Bathonian, and Callovian.

The Lower/Middle Jurassic boundary corresponds to the transition from the cherts to the clastic part of the sequence.

To summarize, the obtained new data enable us to conclude the following:

(1) The age of the cherty allochthons that compose the Samarka prism is established up to the Lower Jurassic inclusive.

Plate 1 Early Jurassic radiolarians from the cherty-clastic sequence of the Samarka accretionary prism. Early Jurassic radiolarians: figs 1, 2 – *Parahsuum simplum* Yao, 1 – Us-21A, 110x, 2 – Us-21, 160x; fig. 3 – *Parahsuum* cf. *simplum* Yao, Us-21, 160x; fig. 4 – *Parahsuum* cf. *magnum* Takemura, Us-23, 80x; fig. 5 – *Mesosaturnalis* sp., Us-21A, 110x; figs 6, 7 – *Hsuum medium* (Takemura), Us-21A, 110x; fig. 8 – *Hsuum medium* (Takemura), Us-23, 110x; fig. 9 – *Parahsuum levicostatum* Takemura, Us-21, 160x; fig. 10 – *Tripocyclia* sp., Us-20, 110x; Middle Jurassic radiolarians: figs 11, 12 – *Hsuum hisuikyoense* Isozaki et Matsuda, Us-13, 11 – 110x, 12 – 160x; figs 13, 14 – *Hsuum robustum* Pessagno et Whalen, Us-13, 160x; fig. 15 – *Hsuum* cf. *belliatulum* Pessagno et Whalen, Us-11, 160x; fig. 16 – *Laxtorum*(?) *jurassicum* Isozaki et Matsuda, Us-10, 110x; fig. 17 – *Parahsuum* cf. *grande* Hori et Yao, Us-13, 160x; fig. 18 – *Parahsuum officerense* Pessagno et Whalen, Us-16, 160x; fig. 19 – *Tricolocapsa* sp., Us-13, 240x; fig. 20 – *Stichocapsa japonica* Yao, Us-13, 160x; fig. 21 – *Andromeda* sp., Us-13, 90x.

(2) The beginning of subduction and, accordingly, development of the accretionary prism is dated as the earliest Middle Jurassic.

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